

The Role of Rhyolites Melts in the Petrogenesis of Tabenken Coal Seam, Northwest Region, Cameroon

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Abstract

In Cameroon, coal has been identified in Bali, Dschang, Fundong, Mamfe and Tabenken. Though its occurrence has been identified, detailed geologic studies in the areas of geochemistry and petrographic have not been done to qualify these coal seams, which knowledge will afford geologists to have a holistic view of coal. The objective of this research thus was carry petrographic and geochemical analyses of the coal samples in other to have an insight about its quality and afford Government and potential investors take informed decisions when considering exploitation and use of the resource. In other to achieve these ten coal samples collected during field studies were subjected to petrographic studies and two samples for geochemical analyses following standard procedures. Results from field description and petrography show that, coal ranks within the seams range from peat to anthracite, some occurring singly, while most were mixed with shale, sandstone and rhyolites. Thin sections show that the coal intercalated with rhyolites has high quartz content. The presence of high quartz percentage may suggest two origins; coal formed from detrital origin and intrusion of silica from rhyolitic tuffs, since rhyolites contain high percentages of silica (>70 %). The presence of hydrocarbon veins observed in the thin sections are interpreted as resulting from the presence of the rhyolitic fluids, whose outpours were caused by diapiric movement, of the coal into the viscous rhyolite. These temperatures thus were sufficiently high to melt the coal into liquids. Results obtained for major elements geochemistry reveal that the percentage composition of the elements SiO₂, Al₂O₃ and Fe₂O₃ are dominant constituents and account for over 57.46 % average weight composition. Other major elements such as K₂O, TiO₂, MgO, Na₂O and MnO are present in low concentrations (average of 3.97 wt %). The high concentration of SiO₂ and Al₂O₃ indicate high detrital quartz and clay mineral content in the coal measures. This is in general agreement with results obtained with work done in the nearby Anambra basin, Nigeria. The high concentrations of SiO₂ (57.46 average wt %), (as observed in thin sections as well) for the samples TBN04, TBN08 and Al₂O₃ content (19.58 average wt %) compared with the concentrations of other major elements in the analyzed samples, suggest a high detrital input from the surrounding of the basin during peat formation. The presence of TiO₂ and P₂O₅ is indicative of basaltic rocks of oceanic environment. These elements also reveal that Si, Al, Ti and K originate mostly from a mixed clay assemblage, which is a constituent of kaolinite and illite. The XRD analytical results show that the TCS is mainly composed of siliceous mineral such as quartz (SiO₂), kaolinite [Al₂(Si₂O₅)(OH)₄]. The mineral suites in the TCS coal samples used are similar with the South African coals. Kaolinite is uniformly distributed in the coal samples.

Keywords: Autonomous Shipping, Autonomous Ship Control System, Formalization of Navigation Operations.

Introduction

Coal has been used worldwide as a source of fuel for household and more especially during the industrial revolution as fuel for the running of engines. At that time several coal mines were opened in England and most of Europe. Moreover, fluctuations in crude oil prices have pushed importing countries into searching for other sources of energy. The use of coal could be an appropriate choice. Although coal can be converted into several products, its conversion into transportation fuels (liquids) in which 47.5 % of crude is also in use and is of prime importance [1].

In Cameroon, coal is a very rare natural resource as research has only found peat in Bali in the North West Region, lignite in Dschang, in Foundong East of Bamenda, Mamfe in the South West Region and unquantified coal seams in Tabenken [2]. Tabenken is situated South of Nkambe town around longitudes ($10^{\circ}47'00''\text{E}$ and $10^{\circ}47'00\text{E}$) and latitudes ($06^{\circ}30'00''\text{N}$ and $06^{\circ}35'00''\text{N}$) in the Donga Mantung Division of the North West Region. Tabenken is underlain by a granitic basement, upon which Tertiary-Quaternary volcanoes uncomfortably lie. The volcanic deposits range from basalt to trachyte and rhyolites [3].

General Geology of Tabenken

Tabenken (Figure 1) is dominantly underlain by granitic rocks as well overlain by volcanic outpours composed of basalts, trachyte, rhyolite and hyaloclastite. While the granites are of the Pan African age, K-Ar age determination show that the volcanoes are Tertiary-Quaternary laid down 31-28 Ma [3].

Tabenken lies in the Nkambe massif of the Oku volcanic field situated North West of the continental segment of the Cameroon

volcanic line. The Cameroon volcanic line CVL (Figure 6) is a major volcano-tectonic feature of West Central Africa. It is a Cenozoic volcanic event characterized by a 1600 km long Y-shaped chain of Tertiary to Recent, generally alkaline volcanoes that follows a trend of crustal weakness and stretches from the Atlantic Island of Pagalu, through the Gulf of Guinea, to the interior of the African continent [3]. It is one of two unique volcanic trends in the world that straddles the continental margin characterized by both oceanic and continental volcanisms. The continental sector enters into mainland of Cameroon from South West to North East of Mount Etinde, Mount Cameroon (that erupted seven times in the last century the most recent being in 2000), Mount Bamboutos, the Oku massif, the Adamawa plateau and the Mandara mountains [4, 5]. Tabenken lies on the Nkambe plateau which is one of the strato volcanoes of the Oku massif.

The coal seam of Tabenken outcrops at full lengths along the Mammie and Kummie stream channels with arkosic sandstones overlain by rhyolite. The rhyolite appears to have cut through or overlain the coal and boiled it to produce hydrocarbons, which subsequently flowed into the fractures of this rhyolite constituting micro veins.

The first research carried out on the Tabenken coal seam was done in 2013 [6]. This work involved ultimate analysis which enabled us to determine the percentage of fixed carbon, hydrogen, nitrogen, sulphur, oxygen and ash content in coal. However, this coal seam was first identified and documented by Ntep et al, 2000. The main objective of this work therefore is to evaluate the geochemistry and petrography of the coal so as to ascertain its proper mineralogical composition considering its close association with igneous rocks.

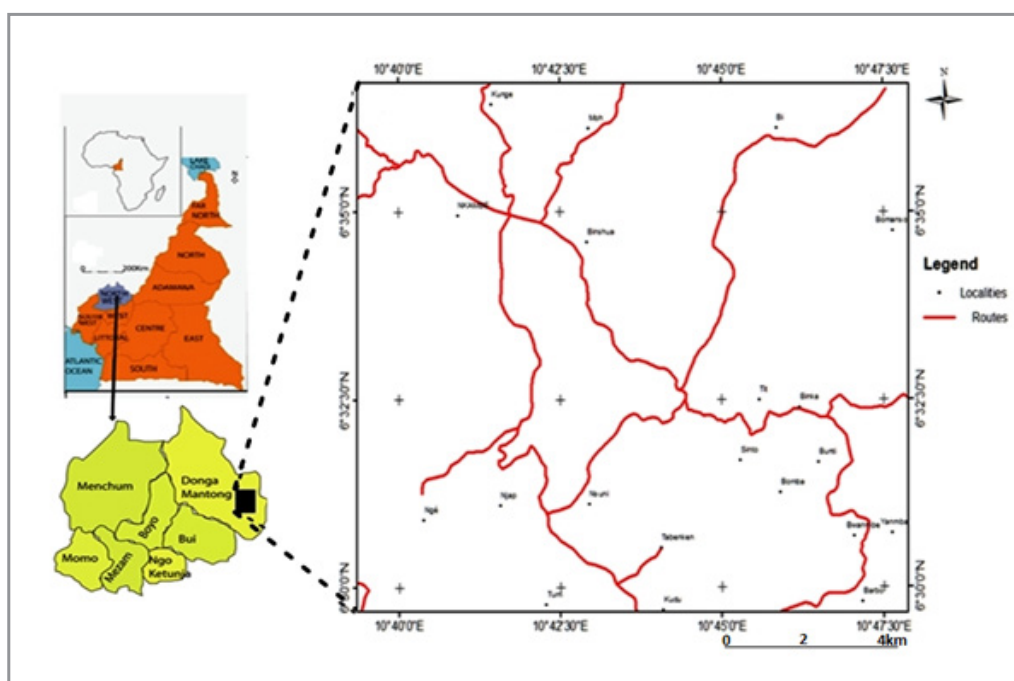


Figure 1: Location map of the study area (We need to increase the font size in the map.)

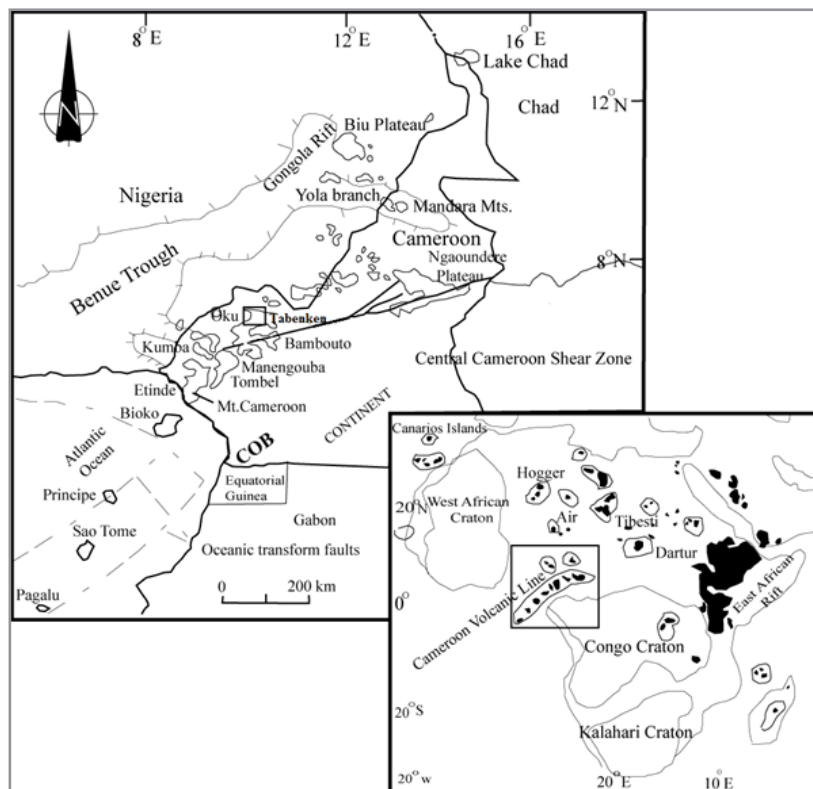


Figure 2: The Cameroon Volcanic Line showing main volcanic centers from Fitton and Dunlop, (1985).

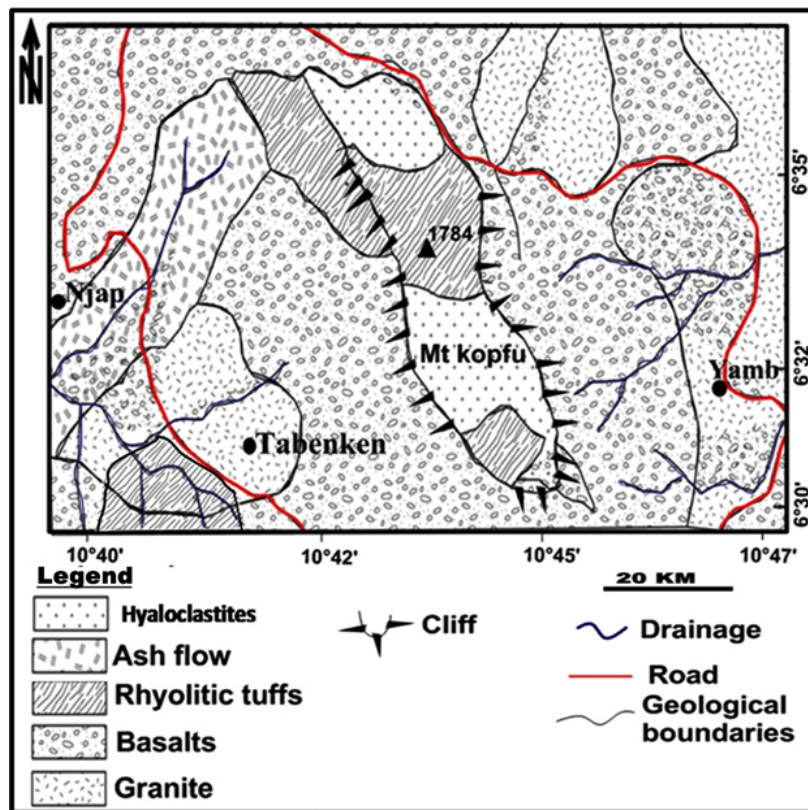


Figure 3: General geology of Tabenken (modified from Njilah, 1991) I suggest we get a section of the map by Tepsogang. This was poorly modified

Identification and Description of the Outcrops

The general outcrop of this area consists of two main formations; igneous formations (granites and volcanics) and sedimentary units as minor outcrops of (clay, conglomerates, arkosic

sandstone and coal).

The Tabenken coal is found along the Kummie and Mammie River valley in Tuku, at longitudes 06°29'32.5"N and latitudes

10°41'16.6E. The exposed basal sequence here consists of conglomerates and arkosic sandstones. These sedimentary formations consist of particles of older rocks which have been transported and deposited. Floats and boulders contain particles of varied sizes and shapes which range from angular to sub angular, rounded to sub rounded and poorly or unsorted pebbles. Downstream, the basal unit consists of shales spread out throughout, with colors ranging from grey reddish brown to pink. Where the rhyolites intrude the coal, the shales grade into units of easily fragmented peat. This unit is overlain by ash beds and patches of mud mixed with sandstone. Further downstream, the basal unit consists of patches of vegetation alongside weathered rhyolites. The rhyolitic tuffs appear with colors ranging from cream white

to pale yellow. They are in direct contact with coal. The rhyolites appear in bands of fragments and are not compacted.

Along the streams, coal is mixed with clay, arkosic sandstone and intercalated with volcanic material. The rounded nature of some of the boulders suggests long distance transportation. There is massive transportation and deposition of material throughout the river bed (Figure 9d) at geographical coordinates N06°29'06.0" and E010°41'21.2", petrified wood can be seen lying along the river bank beneath rhyolites (Figure 9c). This shows a cross section of plant stem with aspects of collenchyma and sclerenchyma clearly displayed.



Figure 4a: Surrounding wall of the upper section of stream bed showing water leakage.

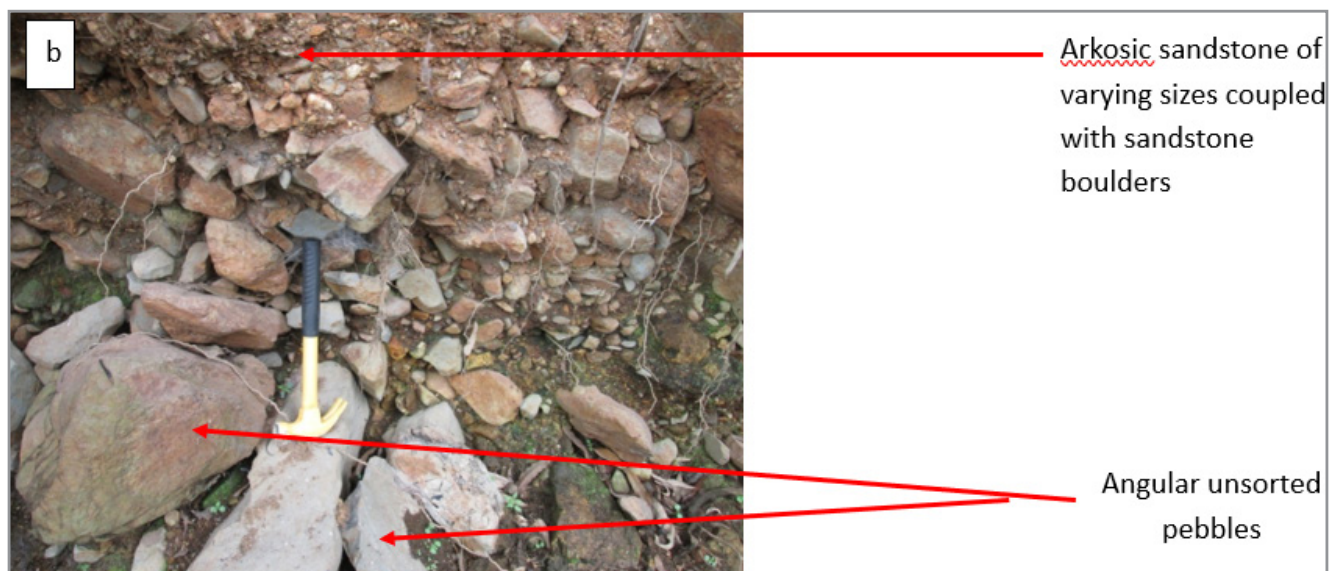


Figure 4b: Conglomerate outlining full length of the mammie river channel.

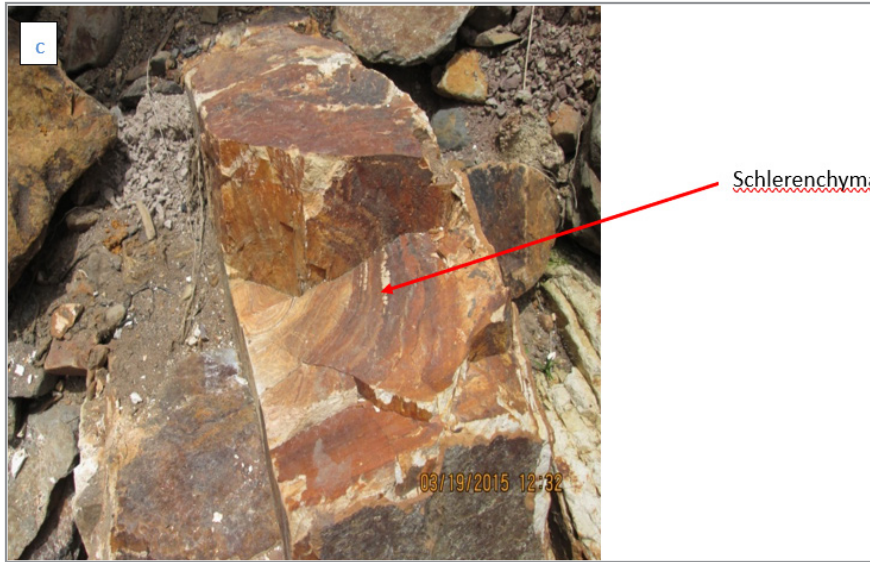


Figure 4c: Petrified wood

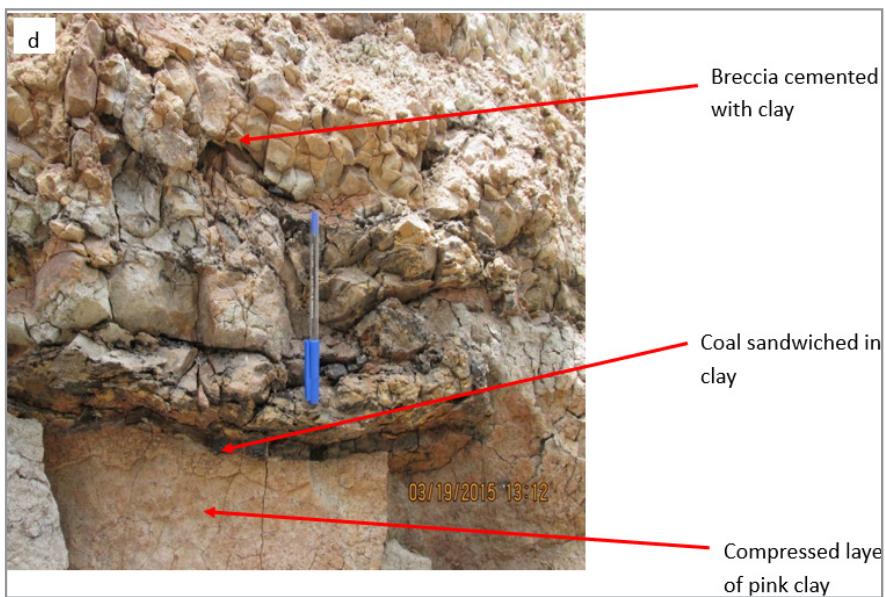


Figure 4d: Coal embedded in clay material

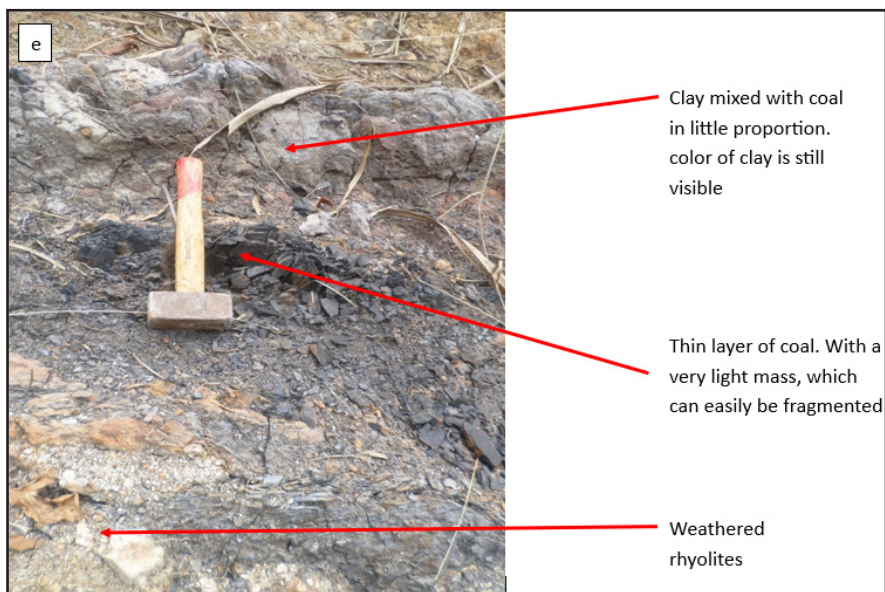


Figure 4e: Coal in its earliest stage of formation (peat).

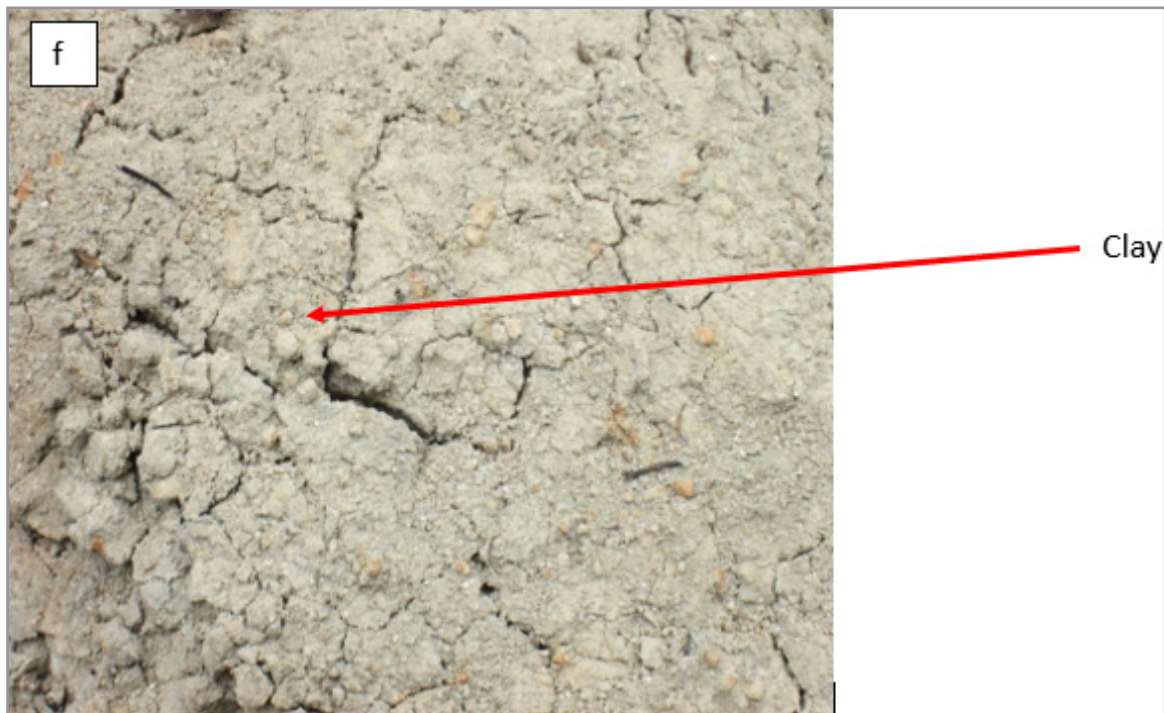


Figure 4f: Clay a typical mineral that spreads throughout the Mammie River channel.

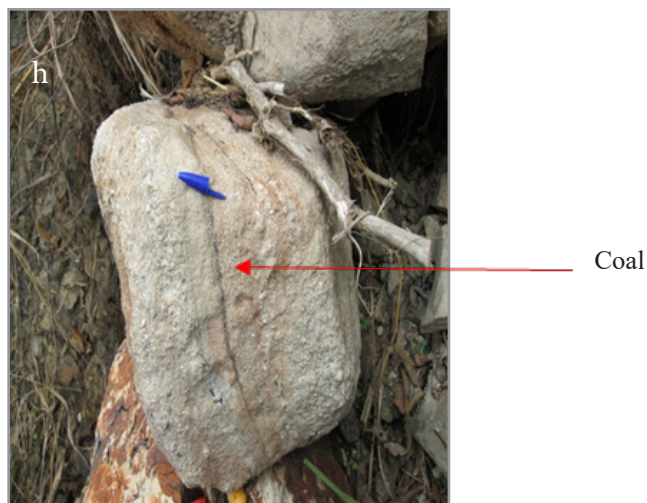


Figure 4: General outcrop covering the Tabenken coal seam:

- g) Coal found within a 6-meter pit dug on the southern portion of the river where cultivation is taking place.
- h) Coal completely cemented and covered with clay and rhyolites.

Field Sampling

A total of 9 samples were collected for geochemical, petrographic and mineralogical analysis. These were labeled with a labeling maker TBN01-TBN09.

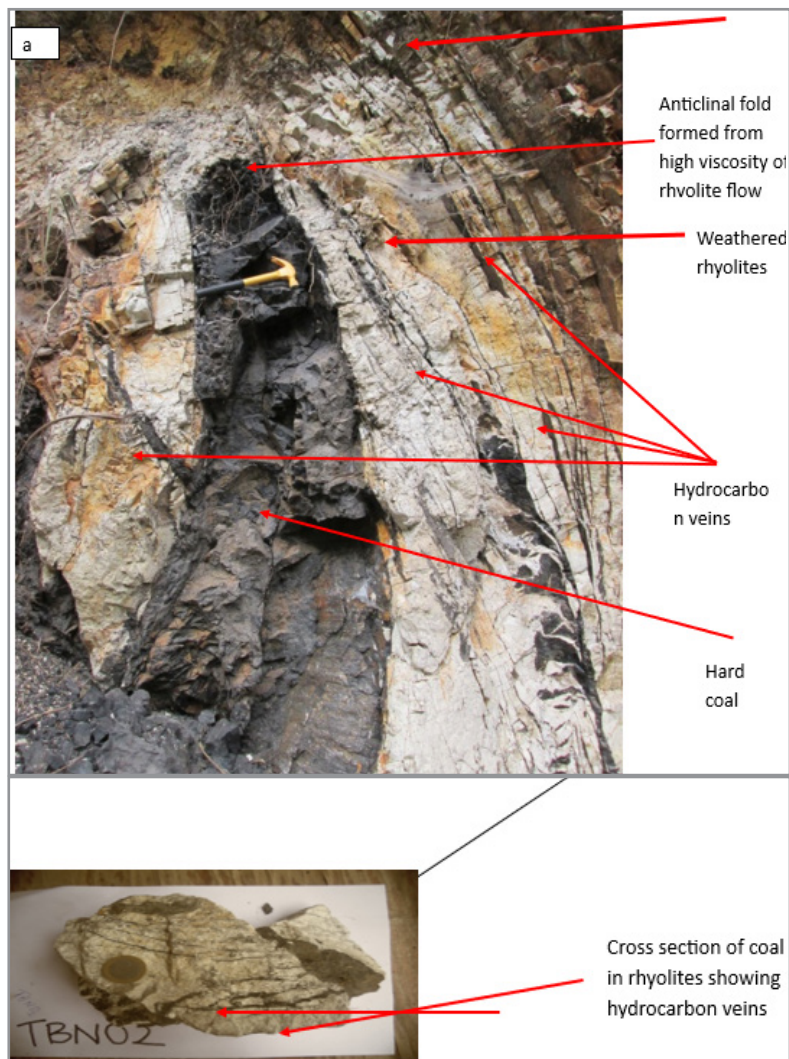








Figure 5: Outcrop showing a coal seam interbedded with sedimentary rocks.

Table 1: Field description of coal and associated rock samples collected from different locations within the study area

Sample name	Location	Description
TBN01	N6°35' and E10°40'	This sample consisted of the different clay types (based on colors observed) that were found around the study area.
TBN02	N06°29'32.4" and E10°41'16.6 	This was the main area of focus for the study. Grey white rhyolite was seen extending a large surface with coal appearing in form of veins. Outcrop was about 6m wide and the coal intercalations appearing in a centimetric direction. Coal layer of about 50 cm thick was seen extending right below the basement and almost extending into the flowing stream. The surface of this formation was covered by vegetation. At this same point we could find thick fine black layer of coal. Coal was folded in an asymmetric direction with the intrusion of yellowish rhyolitic dyke.

<p> TBN03 </p>	<p> N06°29'36.6" and E010°29'36.6" </p> 	<p> Within the river valley at this point coal appeared completely mixed with fine grains, loose and coarse grain sandstone forming a coherent mass. Rock mass was heavy. Grains of quartz could be observed with the eyes. At some points, coal occurred as an interruption in the deposition of sandstone forming a thin continuous layer of coal within the sandstone layer. Sample was cracked with the aid of a geologic hammer. The interior shows a continuous uniform mixture of sandstone and coal with an overall black colour. </p>
<p> TBN04 </p>	<p> N06°29'32.4" and E010°41'16.6" </p> 	<p> Coal was hard with shiny black color and glassy texture. </p>
<p> TBN05 </p>	<p> This was the rhyolite sample I can't lay hand on the picture of this sample </p>	
<p> TBN06 </p>	<p> N06°29'00.1" and E010°41'16.6" </p>	<p> Coal was quite friable, occurring in thin layers which are very fragile and crumble easily into dust. Some where they are coated with reddish brown material. At the base was found black dust and highly fragmented coal of different sizes and variable shapes. </p>
<p> TBN07 </p>	<p> N06°29'30.8" and E010°41'16.6" </p> 	<p> It is fragile woody and fibrous and covered with ground mass. The wood was in its early stage of coal formation and demonstrates characteristics of peat. </p>
<p> TBN08 </p>		<p> This was a piece of coal in sandstone matrix with partially decomposed plant leaves. The color was brownish black. The proportion estimated as 40 % coal and 60 % sandstone. </p>
<p> TBN09 </p>	<p> N06°00'54.41" and E010°41'17.8" </p> 	<p> It was a light weighted coal covered with ash. </p>

Laboratory Analysis Petrography

A total of eight thin sections were prepared at the Institute of

Mining and Geological Research Nkolbisson-Yaoundé from selected samples. The observations of the thin sections were done at the laboratory of Magmatism and Natural hazards, Depart-

ment of Earth Sciences, University of Yaoundé I and photomicrographs were taken using a digital camera.

Geochemistry

Geochemical analysis of major elements of coal and coal bearing rhyolite was carried out on two samples, TBN04 and TBN08, using the ICP-AES and ICP-MS at MIPROMALO.

X-Ray Diffraction Analysis

The X-ray diffraction (XRD) patterns were obtained with a Bruker D8-Avance Eco 1Kw diffractometer (Copper K α radiation, $\lambda=1.5418$ Å, V=40 125 KV, I=25 mA) with Lynxeye Xe energy dispersive detector in the Laboratoire d'Argiles, Géochimie et Environnements Sedimentaires (AGEs) at the University of Liège, Belgium, using standard procedures. The analyses were carried out on the bulk material (non-oriented powder with grinded particles <50 μ m). The XRD patterns were recorded over the 2-70°2 θ angular range for the bulk material. The step sizes considered for this analysis was 0.02°2 θ , whereas the time per step chosen were 0.25 seconds. The primary optic is motorized in order to illuminate a fixed sample length whatever the angular position (16 mm for bulk). The identification of mineral phases was carried out using Eva software.

Data Analysis and Results

Petrographic Description

Macroscopic and microscopic descriptions and the estimation of the modal composition of the minerals were done by observing the various proportions in the thin sections.

The yellowish to greyish white rhyolitic rock TBN02 was found to be fully compacted and partially weathered, with coal appearing in between in form of veins fragments. Observation under the microscope reveals rock texture to be clastic ranging from fine to coarse grains clayey silty to sand. It contains polycrystalline quartz of about 3 mm max and 20 % modal proportion of the whole matrix. It has orthoclase and clay mineral content of 1-2 % modal proportion. The rock matrix is essentially argillaceous coal; (clayey coal) 80-85 % modal proportion.

Coal intercalated in rhyolite (TBN020) in thin section reveals that rock fragment consist of quartz (10-15 %). Their sizes range from semi angular to angular. The rock matrix consists essentially of argillaceous (80-85 %). Rock has a fine clastic to averagely heterogranular texture. It is sandy to clayey.

Sample: TBN02'' (Coal Intercalated in Rhyolite)

Microscopic observation reveals that the rock consists of angu-

lar quartz of 2 mm max occupying (10-15 %) matrix is clayey mixed with coal (90-95 %). The texture ranges from fine clastic to larger grains. Plagioclase has a proportion of <1 % and the ground mass is sandy-clayey. Figure 17G, H.

Sample: TBN08 (Coal Mixed with Sandstone Showing Leafy Portions).

The thin section reveals a clastic texture. The matrix has a modal proportion of (40-45 %) of ferruginous clay. It has micro fractures filled with clay, organic matter and liquids. It has opaque minerals (oxides) which appear angular to sub angular occupying (2-5 %) of the sample. It is dominated with quartz of length (1-0.9 mm) which are angular and sub rounded with a proportion of (50-55 %). Figure 17C, D

Sample: TBN04 (Coal Mixed with Sandstone)

Rock consists mostly of Quartz of modal length 2 mm max under microscopic observation. It has clayey texture with a clay composition of (70-80 %) of the whole rock sample. Microcline 300 μ m max makes up (5-10 %). Matrix is clayey coaly contributing (10-15 %) of rock sample. It is clastic with coarse heterogranular grains.

Sample: TBN05 (Rhyolite)

The rock is made up of phenocrystals, with modal lengths varying between 3- 0.2 mm and modal proportions of (15-20 %) of quartz, feldspath (sanidine), rock fragments (quartzite, granite). Microcline makes up (80-85 %) and has minerals quartz and feldspars (plagioclase, sanidine). It has a microlitic porphyritic texture. Figure 17I, J

- A: TBN02 (coal intercalated in rhyolite observed in plain polarized light),
- B: TBN02' (coal intercalated in rhyolite),
- C: TBN08 (coal mixed with sandstonesandstone observed in polarized light),
- D: TBN08 (coal mixed with rhyolite observed in natural light),
- E: TBN04 (coal in advanced stage of formation observed in plain polarized light),
- F: TBN04 (coal in advanced stage of formation observed in natural light),
- G: TBN02'' (other portions of the same sample of coal intercalated in rhyolite observed in polarized light),
- H: TBN02'' (other portions of sample TBN02 observed in natural light),
- I: TBN05 (rhyolite observed in plain polarized light),
- J: TBN05 (rhyolites observed in natural light).

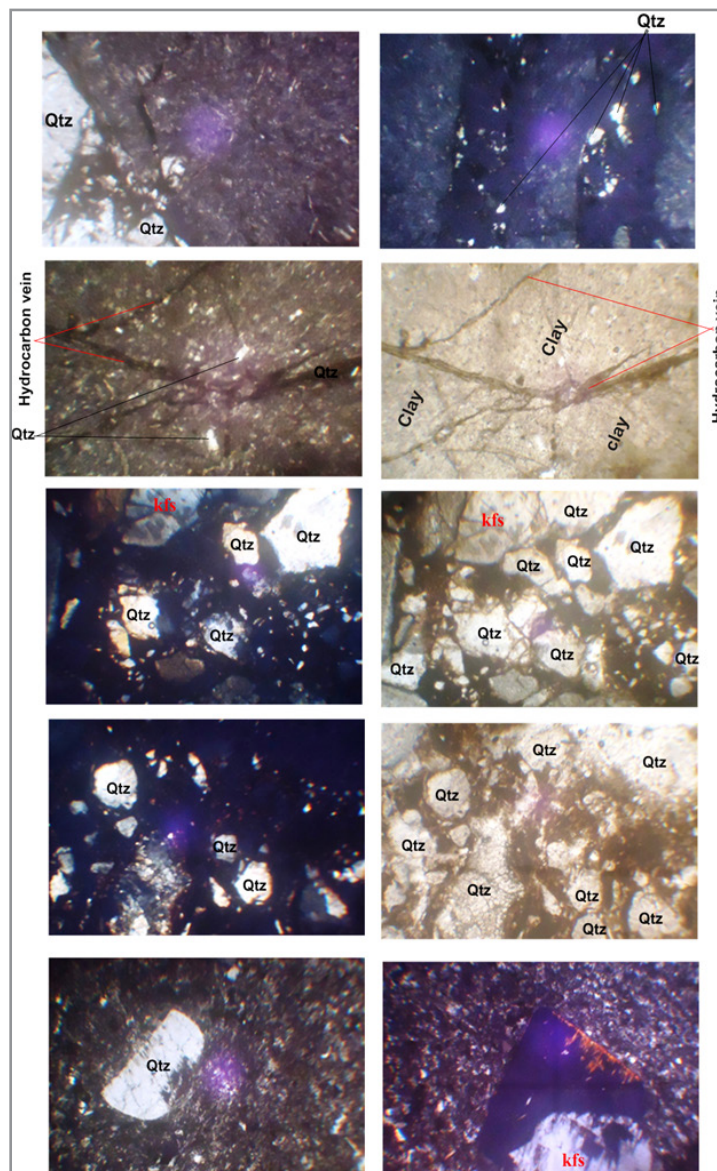


Figure 6: Microscopic features of samples

Major and Trace Element Geochemistry

Major Elements Variation

Major Element's variation in coal: SiO₂ exhibits a wide range within the suite ranging from 49.06 to 65.91 % weight in coal in advanced stage of formation (TBN04) and coal mixed with sandstone (TBN08) respectively. Al₂O₃ decreases sharply in the sample TBN08 17.87 wt% with respect to 21.29 in sample TBN04. Total iron decreases slightly in TBN04-TBN08 2.892-1.816 wt% within the suite, with respect to SiO₂ wt%. CaO ranges from 0.0584-1.266 wt% in the samples TBN08, TBN04 respectively. There is an increase from in K₂O (0.815-1.276)

wt% for TBN04 and TBN08 respectively. MnO is completely absent in sample TBN04, and present in very small amounts 0.0425 % in TBN08. There is a decrease in the percentage of P₂O₅, 0.177 %- 0.297 wt% in TBN08 and TBN04 respectively. MgO is found in coal bearing samples in trace amounts. It increases within the suit from 0.061 %- 0.0943 wt% concentration. Further major element plot following the work (After Suttner and Dutta, 1986) in Yandoka, 2015 of major elements in coal, SiO₂ versus (Al₂O₃ + K₂O + Na₂O) for coal samples TBN04 and TBN08 was done to get a provenance study for the tectonic.

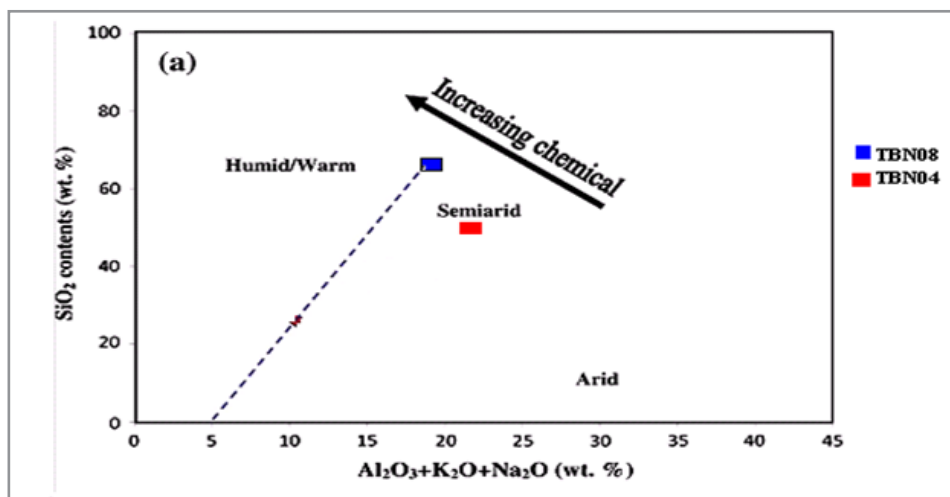


Figure 7: Bivariate plots of SiO₂ versus (Al₂O₃ + K₂O + Na₂O) contents for paleoclimate discrimination (After Suttner and Dutta, 1986).

Table 2: Major element geochemistry of coal bearing samples collected from the field TBN08- coal mixed with sandstonesandstone. TBN04- coal in advanced stage of formation

TBN08		TBN04	
FORMULA	CONCENTRATION %	FORMULA	CONCENTRATION%
SiO ₂	65.91	SiO ₂	49.06
Al ₂ O ₃	17.87	Al ₂ O ₃	21.29
Fe ₂ O ₃	2.982	TiO ₂	2.987
K ₂ O	1.276	Fe ₂ O ₃	1.916
TiO ₂	1.08	CaO	1.266
CaO	0.584	K ₂ O	0.815
MgO	0.51	MgO	0.519
P ₂ O ₅	0.177	P ₂ O ₅	0.297
SO ₃	0.0672	SO ₃	0.175
Na ₂ O	0.061	Na ₂ O	0.0943
ZrO ₂	0.04277	ZrO ₂	0.06604
Mno	0.0425	ZnO	0.0557
SrO	0.0294	Cr ₂ O ₃	0.0434
ZnO	0.0284	CeO ₂	0.041
Cr ₂ O ₃	0.013	V ₂ O ₅	0.0379
CuO	0.005	SrO	0.0214
LOI (actual)	9.31	CuO	0.0136
SUM	99.99	Y ₂ O ₃	0.0124
		NiO	0.0107
		Nb ₂ O ₅	0.0083
		Ga ₂ O ₃	0.005
		Rb ₂ O	0.0036
		LOI (actual)	21.25
		SUM	99.99

Rare Earth Elements (REE) Geochemistry

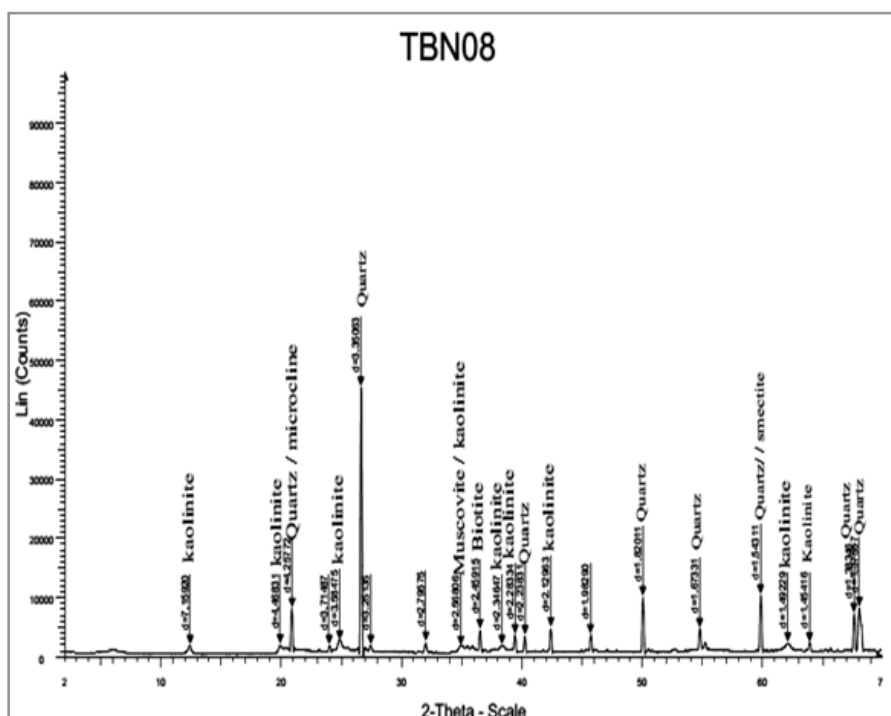
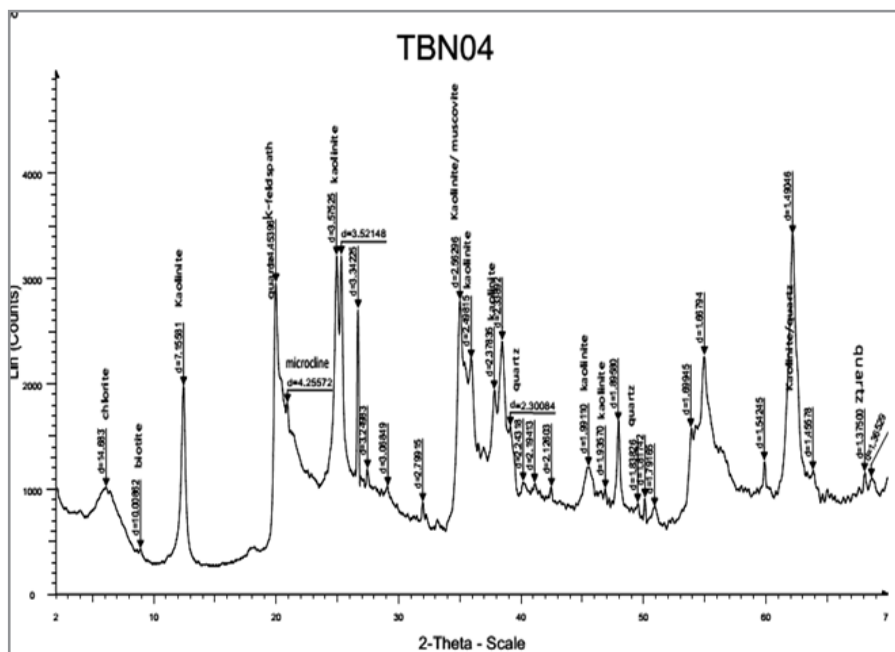
REE elements are a group of elements with atomic number between 57 (La) and 72 (Lu). They are generally characterized by relatively large ionic radii and valences of either +2 or +3. Ionic radius for these elements decreases slightly but steadily as atom-

ic number of REE's increases. REE are generally classified into two groups, LREE, HREE and MREE.

X-Ray Diffraction Analysis

The XRD analytical results (Figure 25) show that the Tabenk-

suites in the coal sample used in the present study are consistent with the previous studies in South African coals [7]. In both samples quartz and kaolinite are the dominant mineral phases



ducted on coal to measure its thermal maturity by vitrinite and huminite reflectances and maceral composition analysis to reconstruct the peat forming conditions. Coal composition is dependent on its origin, degree of carbonization and place of coal basin. From the observations of thin sections, it is seen that, coal intercalated with rhyolites had numerous quartz content. The presence of high quartz percentage may suggest two origins. Firstly, it may suggest coal formed from detrital origin and secondly the intrusion of silica from rhyolitic tuffs since rhyolites contain high percentages of silica above 70 % [8, 9].

The presence of hydrocarbon veins observed in the thin sections are interpreted as resulting from the presence of the rhyolite fluids, whose outpours were caused by diapiric movement, of the coal into the viscous rhyolite. These temperatures thus were sufficiently high to melt the coal into liquids [10].

Interpretation from Coal Geochemistry

Results obtained for major elements geochemistry reveal that the percentage composition of the elements SiO₂, Al₂O₃ and Fe₂O₃ are dominant constituents and account for over 57.46 % average weight composition. Other major elements such as K₂O, TiO₂, MgO, Na₂O and MnO are present in low concentrations (average of 3.97 wt %). The high concentration of SiO₂ and Al₂O₃ indicate high detrital quartz and clay mineral content in the coal measures. This is in general agreement with results obtained with work done in the Anambra basin, Nigeria [12]. The high concentrations of SiO₂ (57.46 average wt %), for the samples TBN04, TBN08 and Al₂O₃ content (19.58 average wt %) compared with the concentrations of other major elements in the analyzed samples, suggest a high detrital input from the surrounding of the basin during peat formation [13].

The presence of TiO₂ and P₂O₅ are indicative of basaltic rocks of oceanic environment. Minerals such as Si, Al, Ti and K are mostly associated with quartz, clay minerals and feldspar. These elements also reveal that Si, Al, Ti and K originate mostly from a mixed clay assemblage, which is a constituent of kaolinite and illite.

SiO₂ is the dominant major element in both samples. SiO₂ is higher in sample TBN08 (coal mixed with sandstone). Constituent of SiO₂ may be probably caused by the inter-bedded sandstone or mudstone [14].

The high percentage of Al in coal is related to the presence of clay minerals. High Al₂O₃ values are explained by the presence of kaolinite and illite minerals observed in the sample. Al is generally enriched in kaolinite [15]. The K/Al ratios in the TCS is relatively low 1.27/17.87 and 0.815/21.29 suggesting the occurrence of kaolinite higher than illite. Such enrichment of kaolinite has been reported in much of the world's coals for example in [16].

Ti originates mostly from clay assemblages constituents with the occurrence of kaolinite and illite. The higher Titanium Ti/Al ratios are as a result of clay mineral input [17]. Ti/Al ratio for the TCS is 1.08/17.87 and 2.987/21.29 respectively for the two samples analyzed. This is relatively high suggesting occurrence of Ti within clay lattice [18, 19].

Ca concentration in TCS sample TBN04 is considerably higher than that of the world average of 1 % weight. This suggests that peat was deposited in mire supplied with calcium rich water [17]. The high Ca concentration is explained by carbonate minerals observed in thin section and mineralogy studies [20].

MgO in coal is related to clay minerals, particularly smectite and chlorite as well as carbonates. Mg concentration in TCS is 0.042 % weight is higher than the world average of 0.02 %. The higher Mg concentration is due to the enrichment of carbonates in TCS [21].

The concentration of K in the oxides of the TCS average weight is (0.01 and 0.009) respectively. This is substantially higher than that of the world's coal average (0.01 weight %), [19]. Higher concentration is mainly attributed to the presence of illite clays in the analyzed coals [22].

Fe in coal occurs in a number of forms such as Fe sulfides (pyrite and marcasite, Fe carbonates, siderite and ankerite) Fe-bearing clay- iron sulphates ferrous sulphate and Jarosite are organically bound Fe [20]. High Fe contents were recorded in the TCS. The iron availability may have been associated with the transport of organic colloids or clay into the basin. (Figure20) Plotting SiO₂ content against Al₂O₃+Na₂O+K₂O reveals that coal formed during semi-arid environmental conditions [23].

Results from x-ray Diffractogram Analysis

Most of the inorganic matter in coal is present as minerals dispersed throughout the coal macerals [24]. The dominant minerals in coals are usually composed of sulfides, clay, carbonates and quartz [25]. Sulfides are composed of pyrite and marcasite but pyrite is dominant by and usually found associated with other sulfides or oxides in coal bed and as a replacement mineral in fossils [26].

The XRD analytical results show that the TCS is mainly composed of siliceous mineral such as quartz (SiO₂), kaolinite [Al₂(SiO₂O₅)(OH)₄]. The mineral suites in the TCS coal samples used are similar with the South African coals [27]. Kaolinite is uniformly distributed in the coal samples. According to Akinyemi, (2012), kaolinite is commonly present in coal as two species with different crystallinity, namely a low crystallinity detrital kaolinite and a high crystallinity neomorphic kaolinite. Kaolinite contains water bound within their lattices [28].

The heat from the rhyolitic intrusions which could have to 1000°C at the time of the injection may have caused changes in the organic matter of the coal. According to Benedict et al., (1968), Teichmuller et al., (1998) this can be changed to natural coke characterized by the presence of a mosaic structure and pores [29]. Anthracites may soften a little if at all, during intrusion and tend to become meta-anthracites, in which case the texture of anthracite is retained, although in chemical composition, optical and other properties, the altered coal approaches graphite, a phenomenon largely observed in Tabenken. The higher rank of some of the samples in Tabenken may be attributed to metamorphism as a result of excessive heat from volcanic activity which then transformed the coal into a higher grade [30]. Such a phenomenon has been observed in coals of Tanjung Enim in the South Sumatra basin in Indonesia [31-81].

Conclusion

The main objectives of this research was carry petrographic and geochemical analyses of the coal samples in order to have an insight about its quality and afford Government and potential investors take informed decisions when considering exploitation and use of the resource. Analyses of field and laboratory data indicated the presence of coal of all grades, ranging from peat to anthracite; coal/sandstone and coal -rhyolite mixtures and as well liquefied coal. The influence of the environment of formation was evident in the geochemistry of the coal with high concentrations of SiO₂ and Al₂O₃ compared to the other elements indicates that the coal is of detrital origin formed during semi-arid conditions.

References

1. Abimnui, N. W. (2013). Geological study of the Tabenken coal seam and its paleo-environmental implications, North West Region, Cameroon (Unpublished master's thesis, 11, 45).
2. Akinyemi, S. A., Gitari, W. M., Akinlua, A., & Petrik, L. F. (2012). Mineralogy and geochemistry of sub-bituminous coal and its combustion products from Mpumalanga Province, South Africa. *InTech*, 47, 57.
3. Kamgang, P., Chazot, G., Njongfang, E., Tchumegnie, N. B., & Tchoua, F. M. (2013). Mantle sources and magma evolution beneath the Cameroon Volcanic Line: Geochemistry of mafic rocks from the Bamenda Mountains (NW Cameroon). *Gondwana Research*, 24, 724–741.
4. Marzoli, A., Peccirillo, E. M., Renne, P. R., Bellieni, G., Lacumin, M., Nyobe, J. B., & Tongwa, A. T. (2000). The Cameroon Volcanic Line revisited: Petrogenesis of continental basaltic magmas from lithospheric and atmospheric mantle sources. *Journal of Petrology*, 41, 87–109.
5. Ngia, R. N., Agyingi, C. M., Foba-Tado, J., Germain, M. M., Nshukwi, A., & Beckley, V. N. (2015). Sedimentology and geochemical evaluation of lignite-argillite sequences in a named basin in Bali Nyonga, Northwest Cameroon. *International Journal of Geosciences*, 917–937.
6. Njilah, I. K. (1991). Geochemistry and petrogenesis of Tertiary–Quaternary volcanic rocks from Mt. Oku Volcano, NW Cameroon (Doctoral dissertation, University of Leeds, 345 pp.).
7. Ntep, G. P., Dupuy, J. J., Matip, O., Fombutu, F. A., & Kalngui, E. (2000). Ressources minérales du Cameroun sur fond géologiques. *MIMEE Cam*, 123–129.
8. Ogala, J. E., Akaegbobi, M. I., Omo-Irabor, O. O., & Robert, B. F. (2009). Statistical analysis of geochemical distribution of major and trace elements of the Maastrichtian coal measures in the Anambra Basin, Nigeria. *Petroleum and Coal*, 51(4), 260–269.
9. Sadeqzadeh, M. B. (2008). Different routes for conversion of coal into liquids: Comprehensive study. In *Proceedings of the World Congress on Engineering and Computer Science (WCECS 2008)*.
10. Say, G. S., & Abdullah, W. H. (2012). Geochemical and petrographical characteristics of low-rank Balingian coal from Sarawak, Malaysia: Its implications on depositional conditions and thermal maturity. *International Journal of Coal Geology*, 96–97, 22–38.
11. Shu, T., Dazhen, T., Hao, X., Jianlong, L., & Xuefeng, S. (2013). Organic geochemistry and elements distribution in Dahuangshan oil shale, southern Junggar Basin: Origin of organic matter and depositional environment. *International Journal of Coal Geology*, 115, 41–51.
12. Singh, A. K., Singh, M. P., & Sen, K. (2003). New classification scheme for microtextures of natural cokes: A case study from heat-affected coking coals of Jharia Coal Field, India. In *Proceedings of the 20th International Pittsburgh Coal Conference*.
13. Hieronymus, B., Kotschoubey, B., & Boulegue, J. (2001). Gallium behavior in some contrasting lateritic profiles from Cameroon and Brazil. *Journal of Geochemical Exploration*, 72, 147–163.
14. Vassilev, S. V., & Vassileva, C. G. (1996). Occurrence, abundance and origin of minerals in coal and coal ashes. *Fuel Processing Technology*, 48, 85–106.
15. Ross, D. J. K., & Bustin, R. M. (2009). Investigating the use of sedimentary geochemical proxies for palaeoenvironment interpretation of thermally mature organic-rich strata: examples from the Devonian-Mississippian shales, Western Canadian Sedimentary Basin. *Chemical Geology*, 260, 1–19.
16. Calvert, S.E., Bustin, R.M., & Ingall, E.D. (1996). Influence of water column anoxia and sediment supply on the burial and preservation of organic carbon in marine shales. *Geochimica et Cosmochimica Acta*, 60, 1577–1593.
17. Stach, E., Mackowsky, M. T., Teichmüller, M., Taylor, G. H., Chandra, D., & Teichmüller, R. (1975). *Stach's Textbook of Coal Petrology*. Gebrüder Borntraeger.
18. Spears, D. A. and Zheng, Y. (1999). Geochemistry and Origin of Elements in Some UK Coals. *International Journal of Coal Geology*, 38, 161–179. [https://doi.org/10.1016/S0166-5162\(98\)00012-3](https://doi.org/10.1016/S0166-5162(98)00012-3)
19. Say, S.G., & Abdullah, W. H. (2011). Concentration and association of minor and trace elements in Mukah coal from Sarawak, Malaysia, with emphasis on the potentially hazardous trace elements. *International Journal of Coal Geology*, 88, 179–193.
20. Wang, J., Yamada, O., Nakazato, T., Zhang, Z. G., Suzuki, Y., & Sakanishi, K. (2008). Statistical analysis of the concentrations of trace elements in a wide diversity of coals and its implications for understanding elemental modes of occurrence. *Fuel*, 87, 2211–2222.
21. Widodo, S., Oschmann, W., Bechtel, A., Sachsenhofer, R.F., Anggayana, K., & Puettmann, W. (2010). Distribution of sulfur and pyrite in coal seams from Kutai Basin (East Kalimantan, Indonesia): implications for paleoenvironmental conditions. *International Journal of Coal Geology*, 81, 151–162.
22. Yandoka, B.M., Sarki, W.H. Abdullah., Abubakar, M.B., Hakimi, M.H., & Adegoke, A. K. (2015). Geochemistry of the Cretaceous coals from Lamja Formation, Yola Sub-basin, Northern Benue Trough, NE Nigeria: Implications for paleoenvironment, paleoclimate and tectonic setting. *Journal of African Earth Sciences*, 56, 70.
23. Snyman, C. P., & Botha, W. J. (1993). Coal in South Africa. *Journal of African Earth Sciences*, 16, 171–180.
24. Amijaya, D. H. (2005). Paleoenvironmental, paleoecological and thermal metamorphism implications on the organic petrography and organic geochemistry of Tertiary Tanjung Enim coal, South Sumatra Basin, Indonesia. [Doctoral dissertation, Aachen University]. Institute of geology and geo-

- chemistry of petroleum and coal.
25. Andrew, C. S. (2002). Coal petrology and the origin of coal macerals: a way ahead. *International Journal of Coal Geology*, 50, 119–134.
 26. Allan, R. W. (2001). Alkaline rocks and carbonates of the world part3. Geological Society.
 27. Andrzej, S., Wasilczyk, A., & Gliwice, S. P. (2012). Petrographic characteristics of hard coals of different degree of carbonification in the aspect of use them for Coal-Water Slurries (CWSS) production.
 28. Albarède, F. (2003). *Geochemistry: An introduction*. Cambridge University Press. ISBN: 0-521-8146-5.
 29. Bagnouls, F., & Gaussen, H. (1957). Les climats biologiques et leur signification. *Ann. Geol.*, 2, 183-220.
 30. Bessoles, B., & Topette, M. (1980). *Geologie de L'Afrique; La chaine panafricaine "zone obile de L'Afrique Centrale (partie sud) et zone Mobile Soudanaise"*. Mem BEG M, Orleans, France.
 31. Barsky, V., Vlasov, G., & Rudnitsky, A. (2009). Composition and structure of coal organic mass. *Analytical Review Chemical Technology*, 3.
 32. Benedict, L.G., Thompson, R.R., Shigo- III, J.J., & Aikman, R.P. (1968). Pseudovitrinite in Appalachian coking coals. *Fuel*, 47, 125-143.
 33. Blander, M. (2011). Calculations of the influence of additives on coal combustion deposits. Argonne National Laboratory.
 34. Burke, K. (2001). "Origin of the Cameroon line of volcano-capped swells." *The Journal of Geology*, 109, 349–362.
 35. COP21. (2016). New research points to falling CO2 emission. Retrieved from <http://www.ff.com/cms>
 36. Cox, K.J., Bell, J D., & Pankhurst, R.J. (1989). *The interpretation of Igneous Rocks*. Unwin Hyman.
 37. Debjani, N., Singh, A.K., & Banerjee, P.K. (2009). Metallurgical Use of Heat Altered Coal: A Case Study. *Journal of Minerals & Materials Characterization & Engineering*. jmmce.org
 38. Diessel, C.F.K. (1992). *Coal-bearing Depositional Systems*. Springer-Verlag.
 39. EIA. (2011). International Energy Annual- Total consumption (thousand short tons). EIA.gov. Retrieved on 2013-05-11.
 40. EIA. (2012). International Energy Statistics coal: Recoverable reserves. Retrieved on 2012-05-31.
 41. Fitton, J.G., & Dunlop, H.M. (1985). The Cameroon line, West Africa and its bearing on the origin of oceanic and continental Alkali Basalts. *Geological Society Special Publication*, 72, 23–38.
 42. Frey, F.A., Green, D. H., & Roy, S.D. (1978). Integrated models of basalts petrogenesis: a study of quartz tholeiites to olivines melilitites from SE Australia utilizing geochemical and experimental petrological data. *Journal of Petrology*, 9, 463–513.
 43. Han, D.X., & Yang, Q. (1980). *Coal geology in china*. Pecking coal industry publishing in Chinese.
 44. Hook, M., & Aleklet, K. (2009). A review on coal to liquid fuels and its coal consumptions. *International Journal of Energy Research*, 33. Wiley Interscience.
 45. Kwiecinska, B., & Petersen, H. I. (2004). "Graphite, Semi-graphite, Natural Coke, and Natural Char Classification-ICCP System." *International Journal of Coal Geology*, 57, 99–116.
 46. Le Bas, M., Le Maitre, R., Streckeisen, A., & Zanettin, B. (1986). A chemical classification of volcanic rocks based on the total alkali-silica diagram. *Journal of Petrology*, 27, 745–750.
 47. Lee, S. (1996). *Alternative fuels*. CRC Press.
 48. Li, Z., Sun, Y., Yu, J., & Liu, D. (2001). Marine transgression events in coal formation from North China Basin. *Energy exploration and exploitation*, 19(6), 559-567.
 49. Lowe, P. A., Schroeder, W. C., & Liccardi, A. L. (1976). Technical economies synfuels and coal energy symposium, solid-phase catalytic coal liquefaction process. American Society of Mechanical Engineers.
 50. Molua, E.L. (2005). The economics of tropical agroforestry farms in Cameroon. *Journal of Forest Policy and Economics*, 7(3), 199-211.
 51. Miyashiro, A. (1978). Nature of alkalic volcanic rock series. *Contributions to Mineralogy and Petrology*, 66, 91-104.
 52. Naah, Y. C. (2015). *Geology, cartography and petrology of Bambili area (North West Region-Cameroon)*. [Unpublished master's thesis, University of Yaoundé 1].
 53. NETL. (2014). Great Plains Synfuels Plant. National Energy Technology Laboratory. Retrieved from www.netl.doe.gov.
 54. Njabo, & Languy. (2000). Birdlife international, important Bird and biodiversity area facts sheet: Njinsing-Tabenken. Retrieved from <http://www.birdlife.org>
 55. Njilah, I. K., Ghogomu, R. T., Bilong, P., & Eno, B. S. (1999). Unsustainable Exploitation of the Kopfu Montane Forest N.W. Cameroon and associated hazards.
 56. Njilah, I. K., Ajonina, H. N., Kamgang, K. V., & Tchindjang, M. (2004). K-Ar, Ages, mineralogy, major and trace element chemistry of the Tertiary-Quaternary lavas from the Ndu volcanic ridge N.W. Cameroon. *African Journal of Science and Technology*, 5(1), 47-56.
 57. Nuwer, R. (2012). A 20 year low in US carbon emission, green energy, the environment and the bottom line. Retrieved from <http://green.blogs.nytimes.com/2012/08/17>
 58. Peronne. (1969). Notice explicative sur la feuille Wum-Banyo. Directions des Mines et de la geologie du Cameroon.
 59. Roser, B. P., & Korsch, R. J. (1986). Determination of tectonic setting of sandstonesandstone mudstone suites using SiO2 content and K2O/Na2O ratio. *Journal of Geology*, 94, 635–650.
 60. Skoog, D. A., West, D.M., & Standley, R. (2014). *Fundamentals of Analytical chemistry*. Belmont: Brooks/Cole, Cengage Learning. ISBN: 0-495-55832-X.Smithsonian. <http://www.Volcano.S.edu/world/volcano.cfm?Vnum=0204-03->. Retrieved 2011-02-1.
 61. Snelling, A., & Mackay, J. (1984). Coal volcanism and Noah's flood. *Ex Nihilo Technical Journal*, 1, 11–29.
 62. Speight, J. G. (2008). *Synthetic fuels handbook: properties, process and performance*. McGraw-Hill Professional.
 63. Speight, J. G. (2005). *Handbook of coal analysis*. John Wiley and Sons Inc.
 64. Stracher, G. B., Anupma, P. E., & Sokol, V. (2011). *Coal and peat fires: a global perspective volume 1- coal geology and combustion*.
 65. Stuart, E. N. (1976). *The origin of coal*. M.S., Acts and Facts.
 66. Stuart, G. W., Fairhead, J. D., Dorbath, L., & Dorbath, C.

- (1985). "Crustal structure of the Adamawa plateau Cameroon." *Rev. Science et Technology*. Retrieved 2011.
67. Stukalova, I. E., & Rusinova, O.V. (2007). "Thermal Alteration of Coal in the Khasyn Coalfield."
 68. Suites using SiO₂ content and K₂O/Na₂O ratio. (1994). *Journal of coal Geology*, 94, 635–650.
 69. Suttner, L.J., & Dutta, P.K. (1986). Alluvial sandstone composition and palaeoclimate. *Framework mineralogy. Journal of Sedimentary Petrology*, 56, 329–345.
 70. Taylor, G. M. (1998). Rank: thermal and tectonic alteration. In *Organic petrology*. Gebruder Borntraeger.
 71. Technology Status Report (TSR). (1999). Coal Liquefaction. Department of Trade and Industry, UK.
 72. Teichmuller, M., Taylor, G.H., & Littke, R. (1998). Oxidized organic matter in coal. In *Organic Petrology*. Gebruder Borntraeger.
 73. Tetsopgang, S., Suzuki, K., & Adacchi, M. (1999). Preliminary CHIME dating of granites from Nkambe area, North-western Cameroon, Africa. *Journal of Earth and Planetary Science. Nagoya University*, 46, 57-70.
 74. Tiwary, R. K. (2001). Environmental impact of coal mining on water regime and its management. *Water, Air, and Soil Pollution*, 132, 185-199.
 75. Valentim, B., Lemos de Sousa, M. J., Abelha, P., Boavida, D., & Gulartlu, I. (2006). Identification of unusual microscopic features in coal and their derived chars: Influence on coal fluidized bed combustion. *International Journal of Coal Geology*, 68(3-4), 202-211.
 76. Watson, E. B., & Harrison, T. M. (1983). Zircon saturation revisited: Temperature and composition effects in a variety of crustal magma types. *Earth and Planetary Science Letters*, 64, 295-304.
 77. Watson, E. B., Harrison, T. M., & Amos, B. A. (2007). Temperature spectra of zircon crystallization in plutonic rocks. *The Geological Society of America*, 635-638.
 78. Weeber, S. L., Cairncross, B., & Falcon, R. M. S. (2000). Mineralogical, petrographic and Geological controls on coal ash fusion temperature from New Clydesdale Colliery, Witbank Coalfield, South Africa. *Journal of the South African Institute of Mining and Metallurgy*, 100(4), 189-195.
 79. Wilkinson, J. F. G., & Le Maître, R. W. (1987). Upper mantle amphiboles and micas with TiO₂, K₂O and P₂O₅ abundances and 100/(g+Fe²⁺) ratios of coon basalts and andesites: implications for modal metasomatism and undepleted mantle compositions. *Journal of Petrology*, 28, 37-73.
 80. White, W. M. (2012). *Geochemistry* (Unpublished manuscript). Retrieved March 14, 2012.
 81. Yang, S., Jiang, S., Zhao, K., & Jiang, Y. (2013). Petrogenesis and tectonic significance of early cretaceous high-Zr rhyolite in the Dazhou uranium district, Gan-Hang Belt, Southeast China. *Journal of Asian Earth Sciences*, 74, 303-315.